# Magnitude and Drivers of the Seasonal Cycle and Interannual Variability of pCO<sub>2</sub> in the Washington Coast OOI Endurance Array

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## Introduction

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- Carbon flux is negative when carbon is drawn into the ocean (carbon sink) and positive when released from the ocean into the atmosphere (carbon source).
- Understanding the drivers that cause the ocean to act as a source or sink is important for future predictions and models regarding climate change and ocean acidification.
- Previous research by Takahashi et al. (2002) suggests that the coastal northwestern U.S. is a carbon sink. However, there is debate about the magnitude and the role of biological and physical drivers.
- This project studies the seasonal cycle and drivers of the partial pressure of CO<sub>2</sub> (pCO<sub>2</sub>) off the Washington coast, a region with little documentation on this topic.

#### Methods

- Data was taken from the Washington Offshore Surface Mooring of the Ocean Observatories Initiative (OOI) Coastal Endurance Array.
- Instruments used: 3-wavelength fluorometer, air-sea pCO<sub>2</sub> instrument, and meteorological sensors.
- Suspect and failed data was flagged by the OOI and removed in Google Colaboratory using the Pandas Data Analysis Library to ensure quality control.

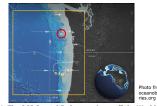


Figure 1. The OOI Coastal Endurance Array off the Washington and Oregon coasts. The Washington Offshore Surface Mooring is denoted by a number 6 and a red circle.

#### Calculations and Analysis

Flux was calculated using the following equation:

**flux** = **k** K<sub>0</sub> (pCO<sub>2,water</sub> - pCO<sub>2,air</sub>) where  $K_0$  is the solubility coefficient and k is piston

velocity from measured wind speed at the mooring.

 Thermal decomposition calculations come from Takahashi et al. (2002):

 $pCO_2 \text{ therm} = \overline{pCO_2} e^{0.0423(T_{obs}-\overline{T})}$  $pCO_2 \text{ nontherm} = pCO_{2,obs} e^{0.0423(\overline{T}-T_{obs})}$ 

where  $pCO_2$  therm is expected  $pCO_2$  with only thermal drivers,  $pCO_2$  nontherm is expected  $pCO_2$  when temperature is held constant, a bar indicates annual mean, and the *obs* subscript indicates observed values.

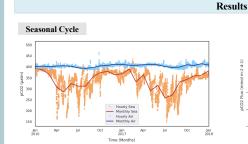
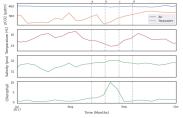
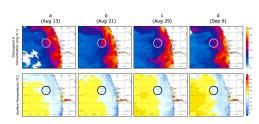


Figure 2. 2016-2017 seasonal cycle and variability of air and seawater pCO<sub>2</sub>. Ocean pCO<sub>2</sub> tends to decrease during late spring and summer, then increases for fail and winter, when it approaches equilibrium with the air. There is more variability of values during late spring and summer, with more constant levels during fail and winter. This seasonal cycle is relatively consistent between the two years; however, the lowest pCo<sub>2</sub> levels are found in late spring for 2016 and mid-summer for 2017.







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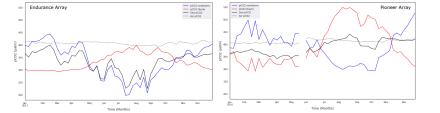
Figure 3. Comparison of weekly pCO<sub>2</sub> flux data from 2016

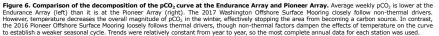
**2018.** Most of the data are below zero, meaning that the area generally acts as a sink. The annual flux value for 2016 was  $-1.9 \pm 2.1$ 

mol m<sup>-2</sup> yr<sup>-1</sup>, and -2.1  $\pm$  1.9 mol m<sup>-2</sup> yr<sup>-1</sup> for 2017 (negative values are flux into the ocean). However, the region can act as a source in late

winter and early spring. There is a high flux value between February and March of 2018, which could be due to an upwelling event.

Figure 4. 2017 weekly data showing a low pCO<sub>2</sub> event in late August. Chlorophyll levels spiked, while pCO<sub>2</sub> levels decreased. Salinity increased slightly and temperature decreased, which would be consistent with coastal upwelling. Dashed gray lines and associated letters denote the dates used for satellite data (see Fig. 5). Figure 5. Satellite images of a low pCO<sub>2</sub> event in 2017. Images are from NASA Giovanni MODIS 8-daily average. Cricics denote the location of the mooring. Letters correspond to those in Fig. 4. The increase in chlorophyll and decrease in temperature at letter 'b' are consistent with upwelling along the coast, after which cold, chlorophyll-rich water is advected toward the Offshore Mooring. These effects are still seen in 'c' and 'd' along the coast, but the water at the mooring becomes warmer and less productive in the last two sets of images.





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### Conclusions

- The region is a net carbon sink, but can occasionally become a carbon source in the winter. The annual flux value for 2016 is -1.9 ± 2.1 mol m<sup>-2</sup> yr<sup>-1</sup>, and -2.1 ± 1.9 mol m<sup>-2</sup> yr<sup>-1</sup> for 2017.
- The  $pCO_2$  values for 2016 and 2017 range from 148  $\mu atm$  to 436  $\mu atm.$
- The seasonal cycle is strongly driven by nonthermal factors, such as upwelling and algal blooms.
- Temperature works to decrease the magnitude of pCO<sub>2</sub> during the winter, keeping the region from becoming a carbon source.
- Evans et al. (2011) estimated annual flux off the Oregon shelf to be -0.3 ± 6.8 mol m<sup>-2</sup> yr<sup>-1</sup>, suggesting that the Washington shelf is potentially a larger sink despite the close proximity of the stations. This indicates that this study would likely not yield a completely accurate estimate of the entire northwestern U.S. coastal region.
- The drivers of the seasonal pCO<sub>2</sub> cycle differ from the Endurance Array to the Pioneer Array. The Washington Mooring is most influenced by nonthermal factors, while the Pioneer Mooring weakly follows thermal factors.
- Despite differences in drivers, the Pioneer Offshore Surface Mooring has a similar annual flux value of  $-1.2 \pm 2.2$  mol m<sup>-2</sup> yr<sup>-1</sup>.

#### **References and Acknowledgements**

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